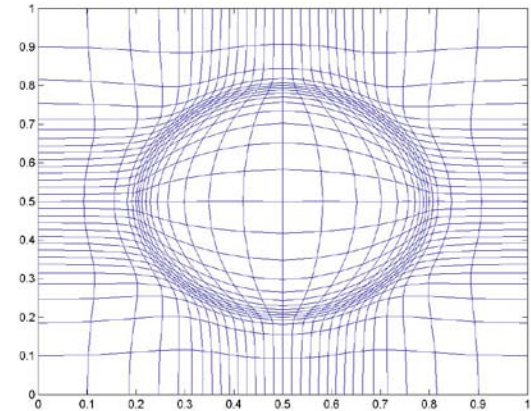
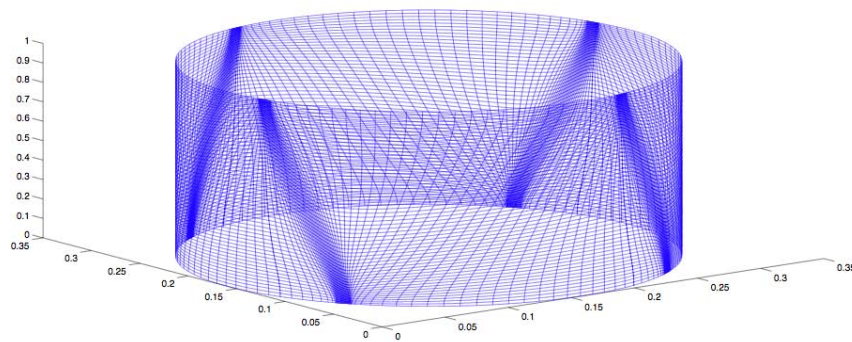


Geometry Based Numerical Methods for Weather Forecasting

Chris Budd (Bath),
Emily Walsh (Bath, SFU), Phil Browne (Bath)

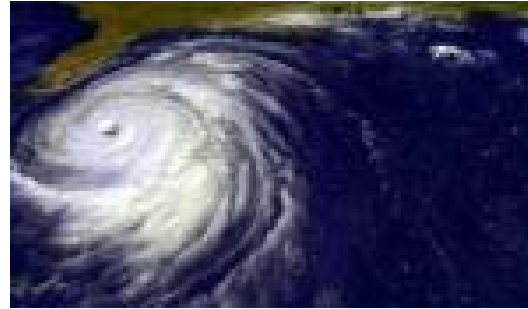


Bath Institute for Complex Systems



UNIVERSITY OF
BATH

Weather forecasting is an important application of mathematics



Forecasts are typically made solve this using a numerical method based on a **computational mesh**

Often need to **locally refine a mesh to capture small scale features**

- (i) For accurate numerical computation eg. Storms
- (ii) For accurate assimilation of observed data

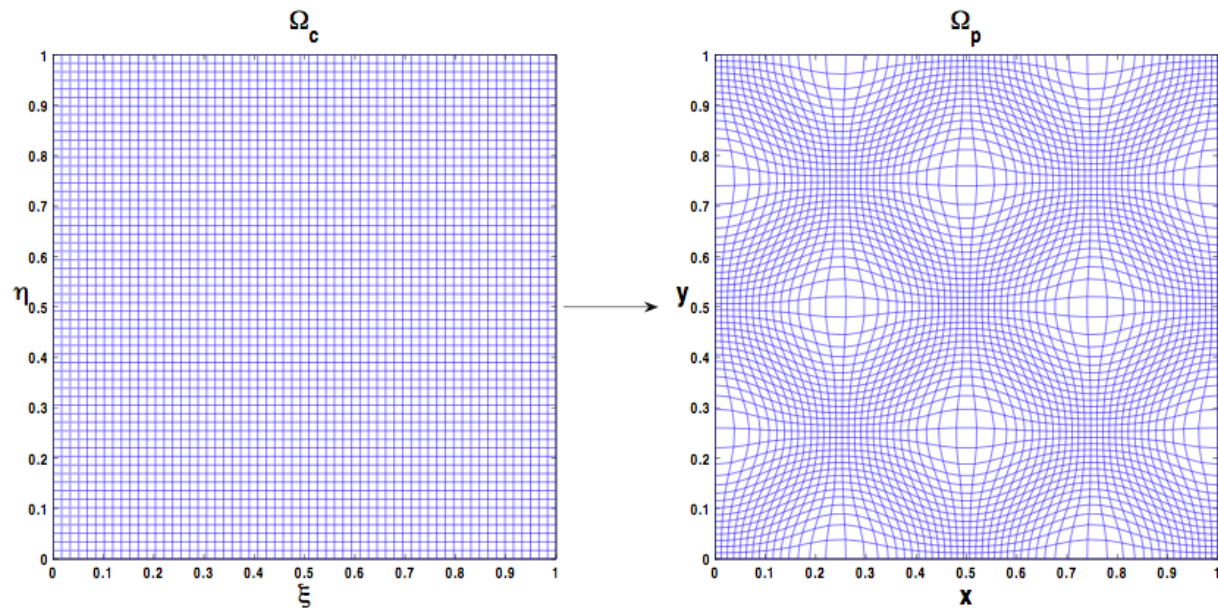
Talk will describe a method for doing this based upon geometrical ideas: **optimal transport**

Geometrical strategy

Have a **computational domain** $\Omega_C(\xi, \eta, \zeta)$

Physical domain $\Omega_P(x, y, z)$

Identify a **map** $F : \Omega_C(\xi, \eta, \zeta) \rightarrow \Omega_P(x, y, z)$



Map a regular mesh, to a mesh used for computation



Mesh used to compute a 3D weather front

Determine F by Equidistribution

Introduce a positive **unit measure** $M(x,y,z,t)$ in the physical domain which controls the mesh density

A : set in computational domain



$F(A)$: image set



Equidistribute image with respect to the measure

$$\int_A d\xi d\eta d\zeta = \int_{F(A)} M(x,y,z,t) dx dy dz$$

Differentiate to give:

$$M(x, y, z, t) \frac{\partial(x, y, z)}{\partial(\xi, \eta, \varsigma)} = 1$$

Basic, nonlinear, equidistribution mesh equation

Choose M large to concentrate points where needed without depleting points elsewhere

Note: All meshes equidistribute some function M

[Radon-Nikodym]

Choice of the monitor function $M(X)$

- Physical reasoning

eg. Vorticity, arc-length, curvature,

- A-priori mathematical arguments

eg. Scaling, symmetry, simple error estimates

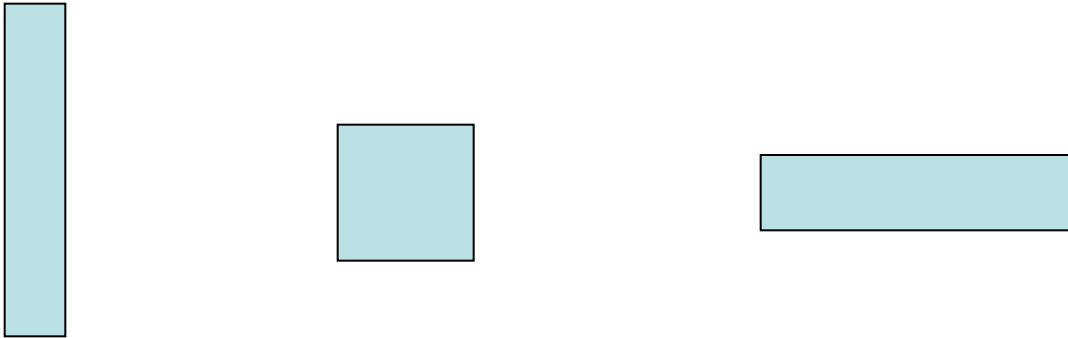
- A-posteriori error estimates

eg. Residuals, super-convergence

- Data correlation estimates

Mesh construction

Problem: in two/three -dimensions equidistribution does **NOT** uniquely define a mesh!

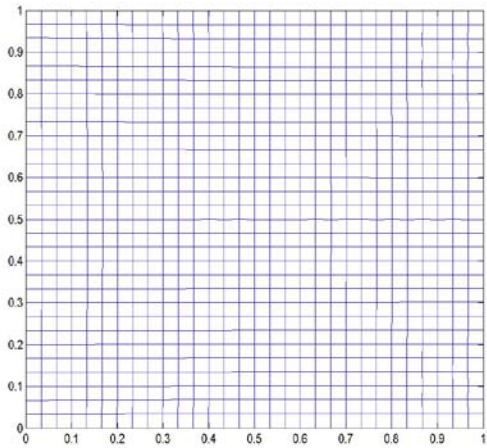


All have the **same area**

Need **additional conditions** to define the mesh uniquely

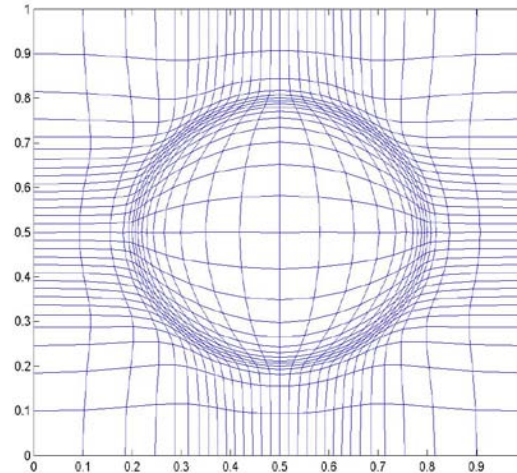
Also want to avoid **mesh tangling** and **long thin regions**

Optimally transported meshes



$$\Omega_C(\xi, \eta)$$

F



$$\Omega_P(x, y)$$

Argue: A good mesh for solving a pde is often one which is as close as possible to a uniform mesh

Monge-Kantorovich optimal transport problem

Minimise $I(x, y, z) = \int_{\Omega_c} |(x, y, z) - (\xi, \eta, \zeta)|^2 d\xi d\eta d\zeta$

Subject to $M(x, y, z, t) \frac{\partial(x, y, z)}{\partial(\xi, \eta, \zeta)} = 1$

Also used in image registration, meteorology

Optimal transport helps to prevent small angles, reduce mesh skewness and prevent mesh tangling.

Key result which makes everything work!!!!

Theorem: [Brenier]

(a) There **exists** a **unique** optimally transported mesh

(b) For such a mesh the map **F** is the **gradient** of a
convex function $P(\xi, \eta, \varsigma)$

P : Scalar **mesh potential**

Map **F** is a **Legendre Transformation**

Some 2D corollaries of the Polar Factorisation Theorem

$$(x, y) = \nabla_{\xi} P = (P_{\xi}, P_{\eta})$$

Gradient map

$$x_{\eta} = y_{\xi}$$

Irrotational mesh

Same construction works in all dimensions

It follows immediately in 2D that

$$\frac{\partial(x, y)}{\partial(\xi, \eta)} = H(P) = \det \begin{pmatrix} P_{\xi\xi} & P_{\xi\eta} \\ P_{\xi\eta} & P_{\eta\eta} \end{pmatrix} = P_{\xi\xi} P_{\eta\eta} - P_{\xi\eta}^2$$

Hence the **mesh equidistribution** equation becomes

$$M(\nabla P, t) H(P) = 1 \quad (\text{MA})$$

Monge-Ampere equation: fully nonlinear elliptic PDE

Properties of the mesh can be deduced from the **regularity of the solution** of the MA equation

Basic idea: Solve (MA) for P with appropriate (Neumann or Periodic) boundary conditions

Good news: Equation has a unique solution

Bad news: Equation is very hard to solve



Good news: We don't need to solve it exactly, and can instead use parabolic relaxation

Q → P

Alternatively: Use Newton [Chacon et. al.]

Use a variational approach [van Lent]

Relaxation in n Dimensions

$$\varepsilon(I - \alpha \Delta_{\xi}) Q_t = \left(\overline{M}(\nabla Q) H(Q) \right)^{1/n} \quad (\text{PMA})$$

Spatial smoothing
[Hou]

(Invert operator
using a spectral
method)

Averaged
monitor

Ensures right-
hand-side scales
like Q in nD to give
global existence

Parabolic Monge-Ampere equation

Solution Procedure

If M is prescribed then the PMA equation can be discretised in the computational domain and solved using a forward Euler method (this is a fast procedure)

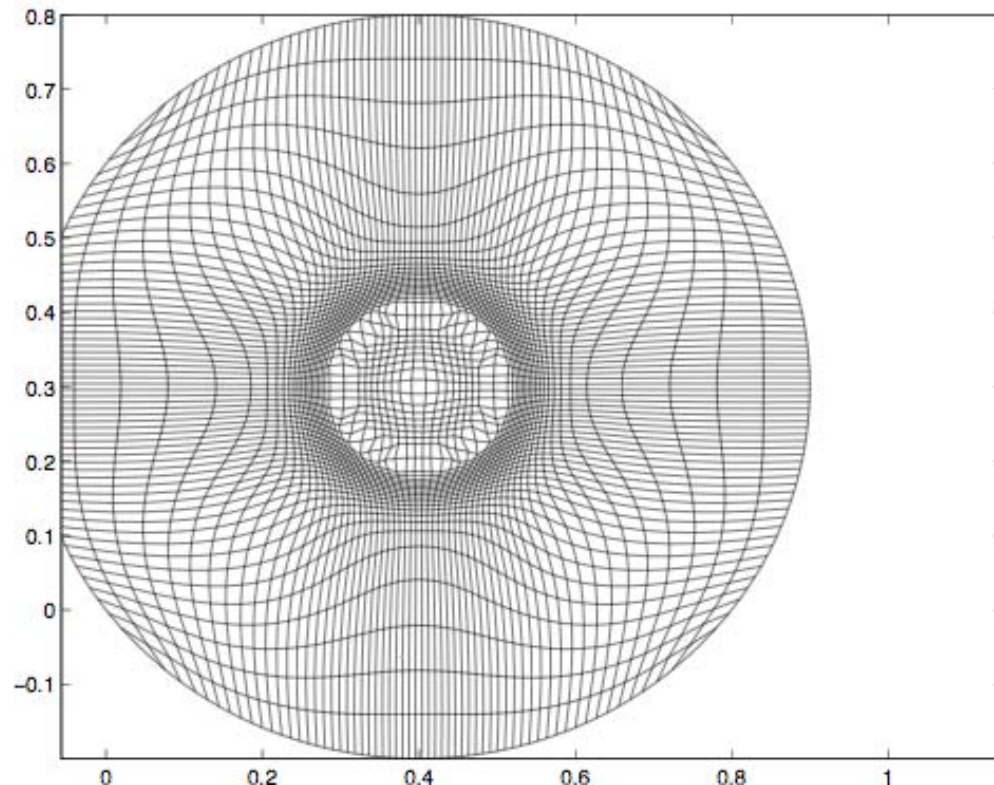
Applications

- Image processing
- Meteorological Data assimilation:

Take M to be the Potential Vorticity of the 3D flow

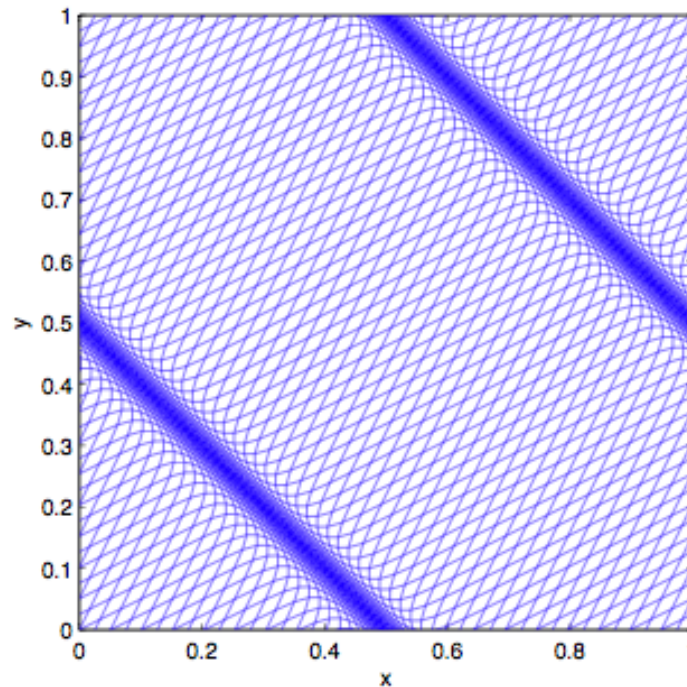
Because PMA is based on a **geometric approach**, it has **a set of useful regularity properties**

1. System invariant under **translations**, **rotations**, **periodicity**



Lemma 1: CJB, EJW [2012]

The solutions of the MA equation exactly align with global linear features



Alignment follows from a close coupling between the local structure of the solution and the global structure. This is NOT a property of other mesh generation methods

2. Convergence properties of PMA

Lemma 2: [Budd, Williams 2006]

(a) If $M(x,t) = M(x)$ then PMA admits the solution

$$Q(\xi, t) = P(\xi) + \Lambda t$$

$$x(\xi) = \nabla_{\xi} Q = \nabla_{\xi} P$$

(b) This solution is **locally stable/convergent** and the **mesh evolves to an equidistributed state**

Proof: Follows from the **convexity** of P which ensures that PMA behaves locally like the **heat equation**

This result is important when initializing a mesh to the initial data for a PDE

Lemma 3: [B,W 2006]

If $M(x,t)$ is slowly varying then the grid given by PMA is epsilon close to that given by solving the Monge Ampere equation.

Lemma 4: [B,W 2006]

The mapping is 1-1 and convex for all times:

No mesh tangling or points crossing the boundary

4. For **appropriate choices of M** the coupled system is **scale-invariant**

Lemma 5: [B,W 2005] **Multi-scale property**

If the PDE has certain continuous group invariants then meshes can be constructed with the same invariance.

This leads to discrete **Noether type** theorems

Lemma 6: [B,W 2009] **Self-similarity**

Such constructions can admit discrete self-similar solutions

Extremely useful properties when working with PDEs which have **natural scaling laws**

Coupling to a PDE

More usually M is a function of the solution of a PDE

- **Carefully** discretise PDE & PMA in the **computational domain**

QuickTime™ and a
decompressor
are needed to see this picture.

Solve the **coupled mesh and PDE system** either

Method One

As **one large system** (stiff!)

Velocity based Lagrangian approach. **Works well for parabolic blow-up type problems (JFW)**

Advantages:

No need for interpolation

Mesh and solution become one large dynamical system and can be studied as such

Disadvantage: Equations are very hard to solve especially when the PDE is strongly advective

Method 2

By alternating between PDE and mesh

1. Time march the PDE
2. Construct a new mesh
3. Interpolate solution onto the new mesh
4. Repeat from 1.

Advantages:

Very flexible, can build in conservation laws

Disadvantage: Interpolation is difficult and expensive

Example 1: Parabolic blow-up

$$u_t = u_{xx} + u_{yy} + u^3, \quad u \rightarrow \infty \quad t \rightarrow T$$

Length scale: $L(t) = (T - t)^{1/2} |\log(T - t)|^{1/2}$

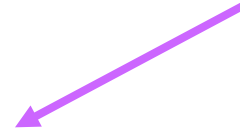
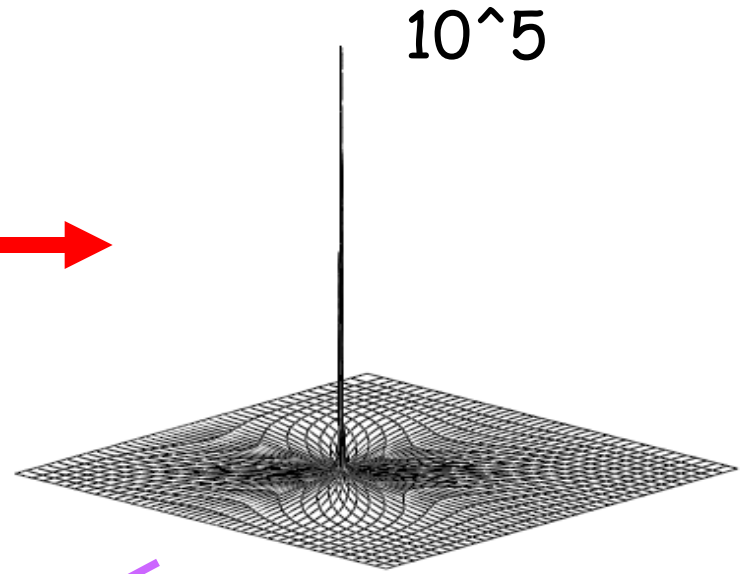
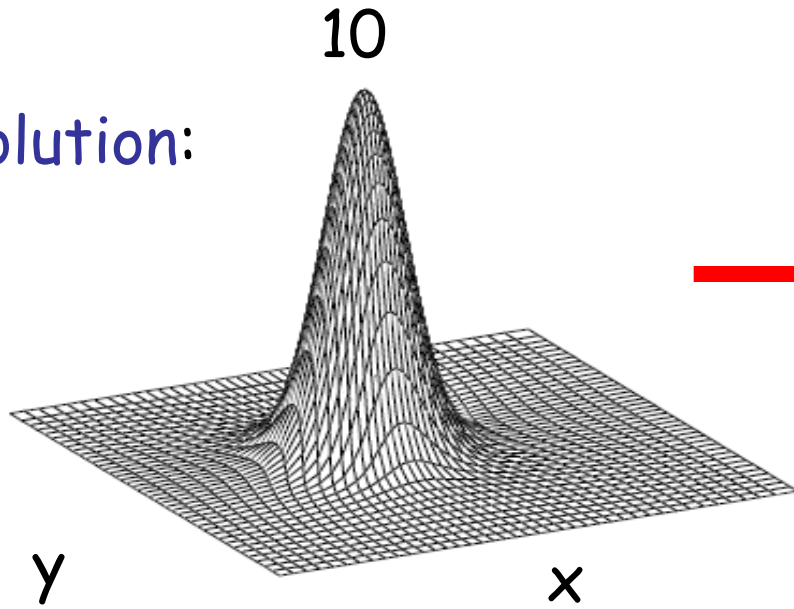
$$M(x, y, t) = \frac{1}{2} \frac{u(x, y)^4}{\int u^4 dx dy} + \frac{1}{2}$$

M is locally scale-invariant, concentrates points in the peak and keeps 50% of the points away from the peak

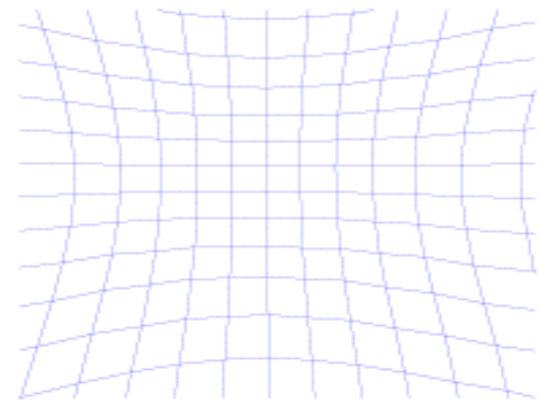
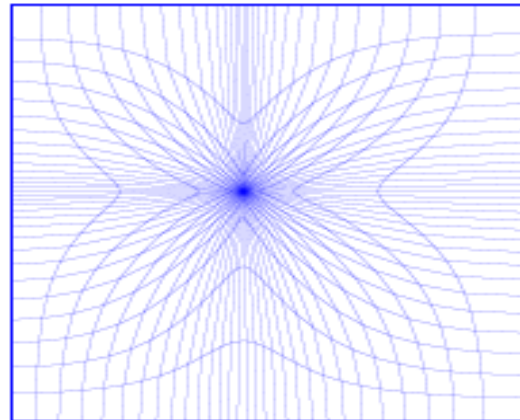
Solve using Method 1

$$u_t = u_{xx} + u_{yy} + u^3$$

Solution:

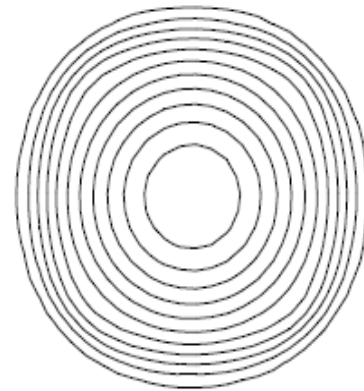
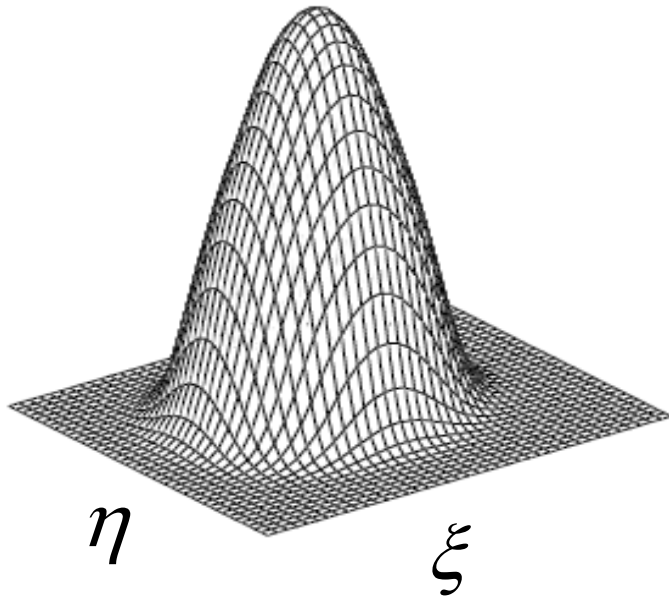


Mesh:

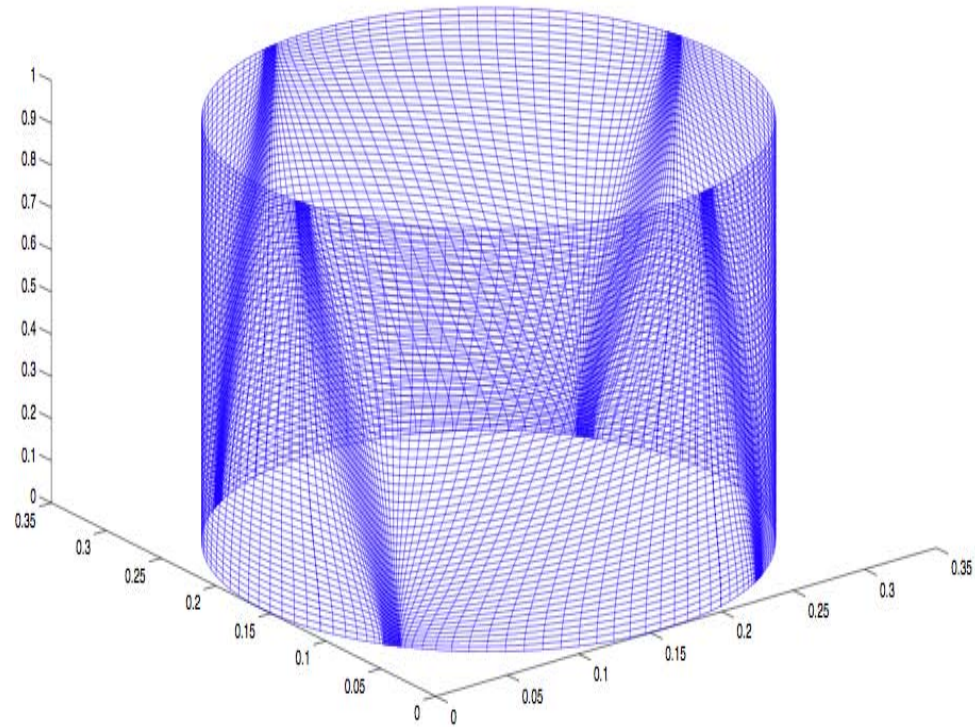
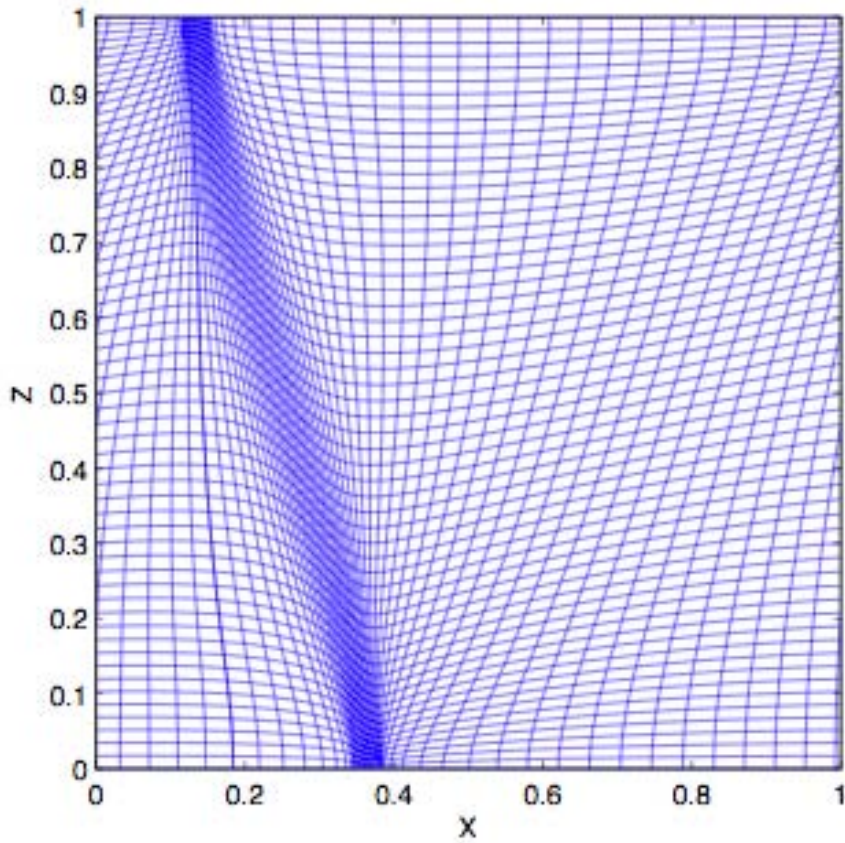


Solution in the computational domain

10^5



Example 2: Tropical storm formation (Eady problem)



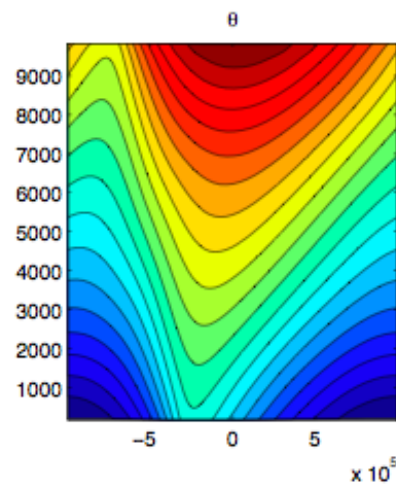
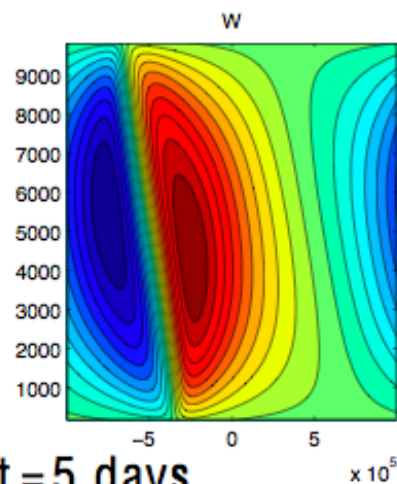
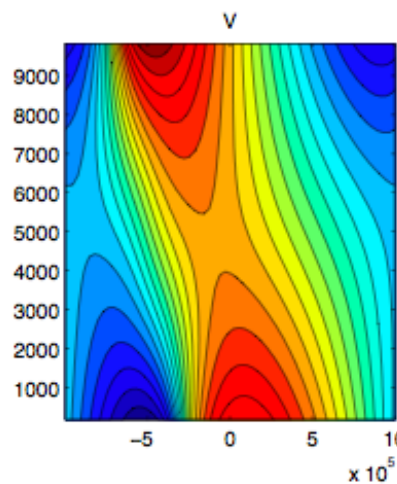
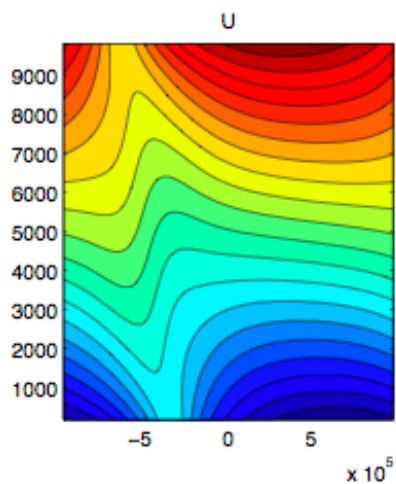
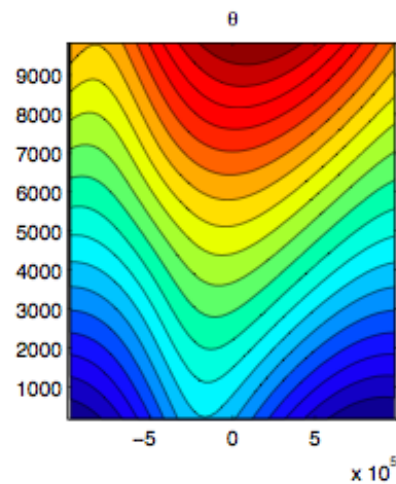
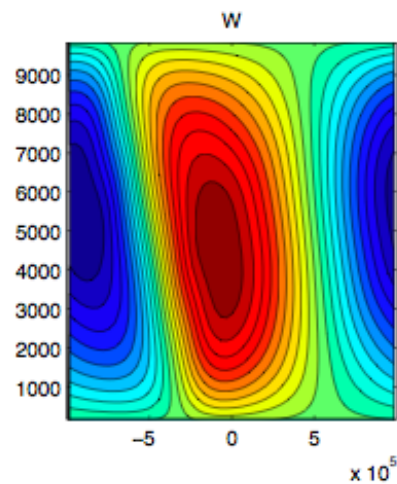
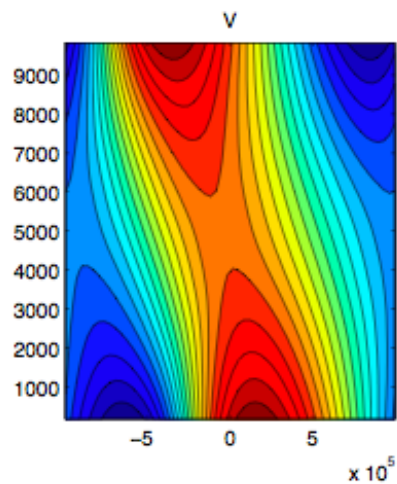
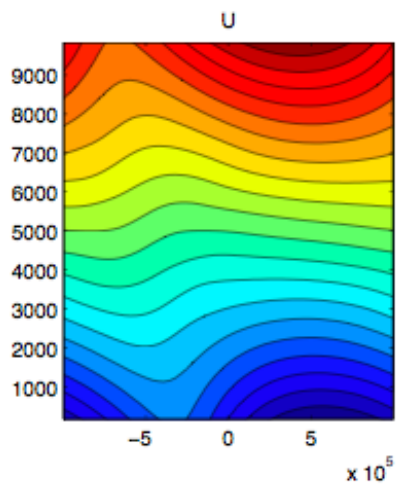
QuickTime™ and a
decompressor
are needed to see this picture.

M : Maximum eigenvalue of Potential Vorticity R

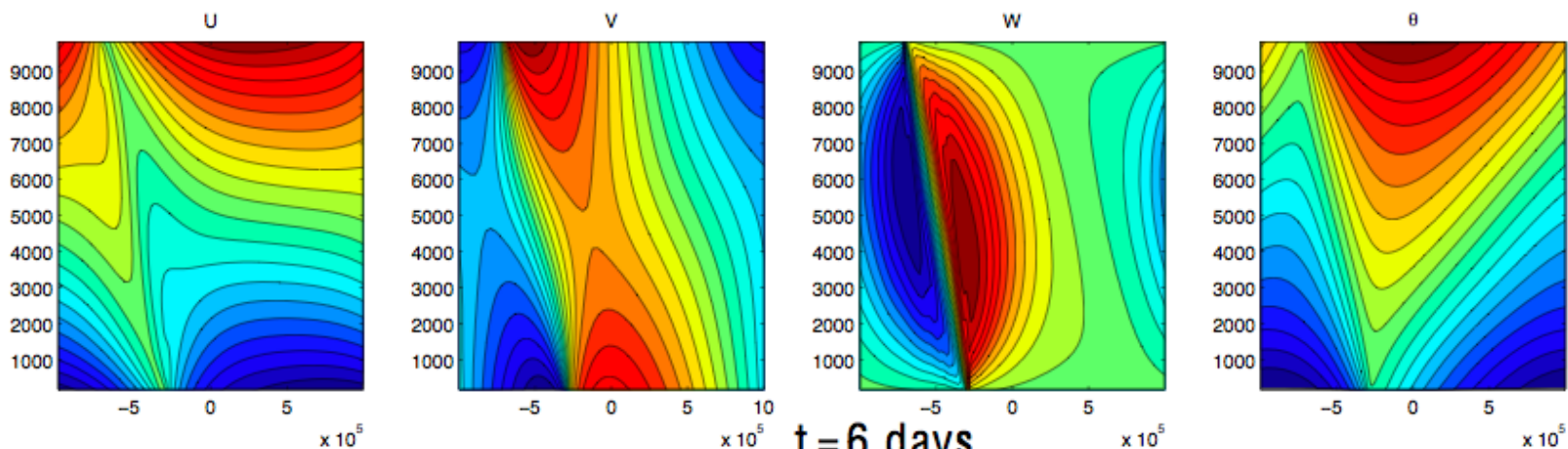
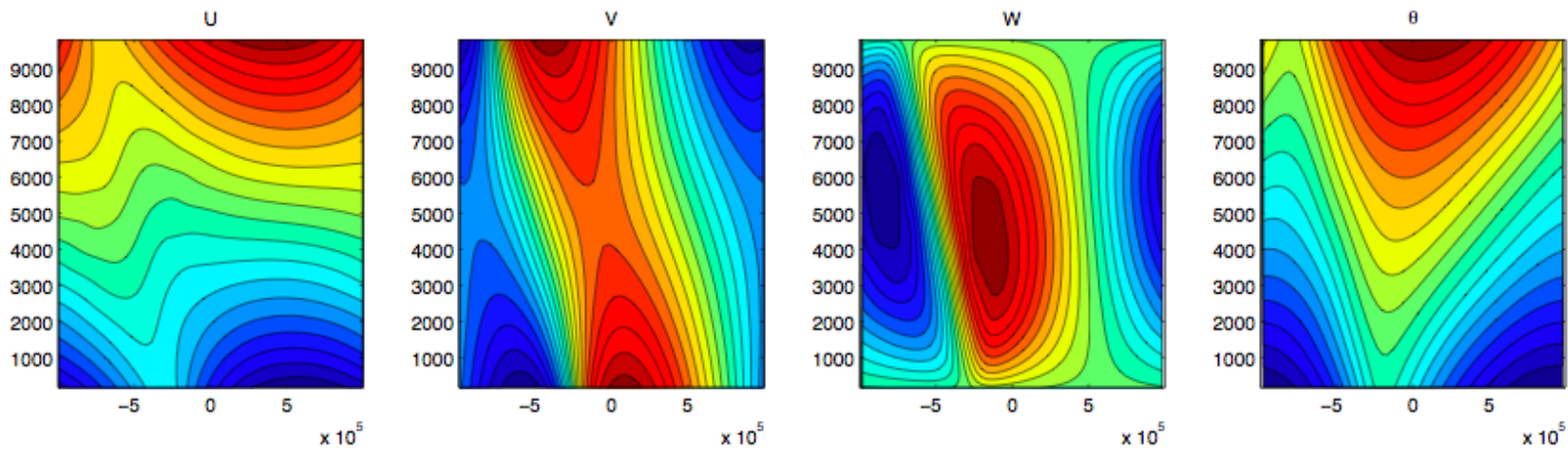
$$R = \begin{pmatrix} f^2 + f v_x & f v_z \\ g \theta_0^{-1} \theta_x & g \theta_0 \theta_z \end{pmatrix}$$

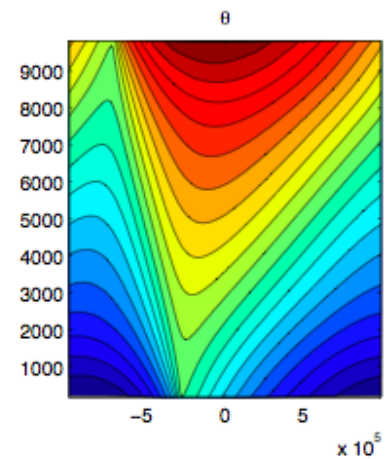
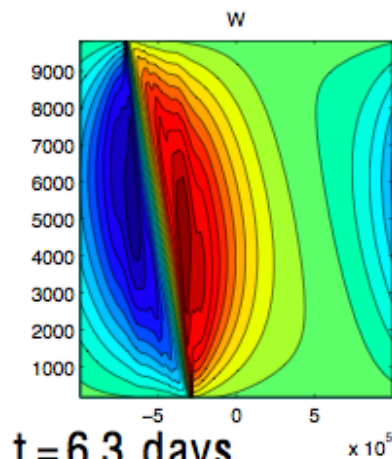
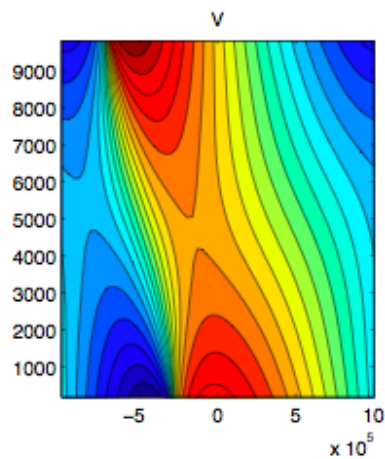
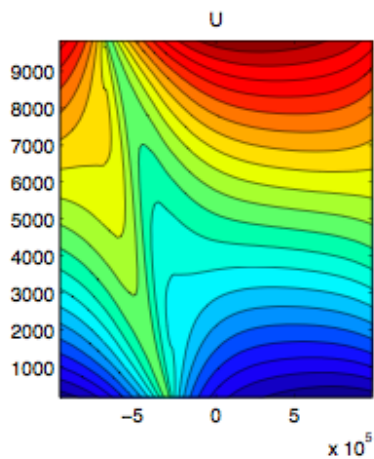
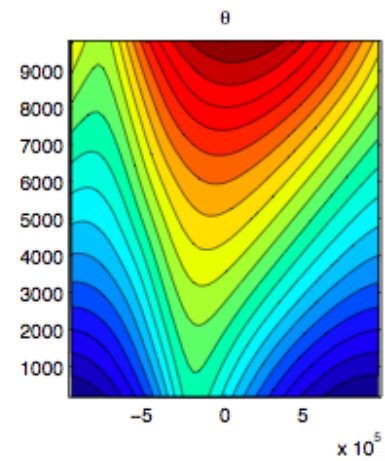
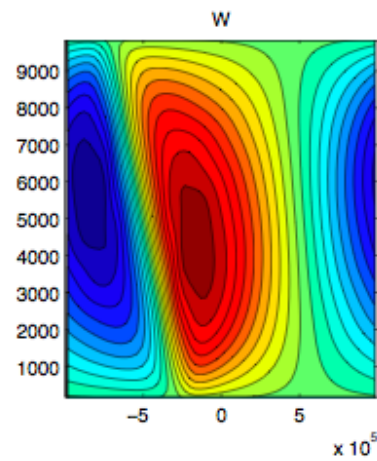
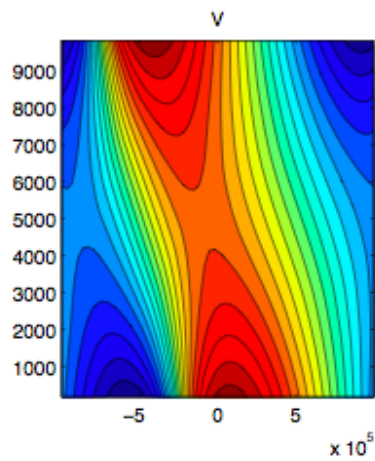
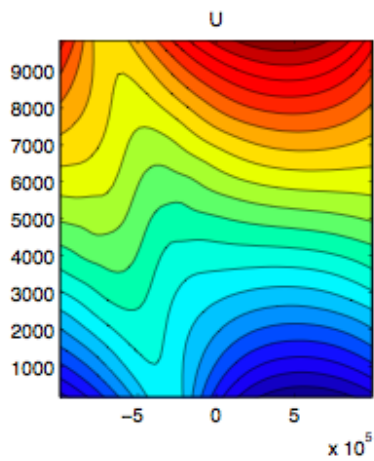
Solve using Method 2, with pressure correction

- Update solution every 10 mins
- Update mesh every hour
- Advection and pressure correction on adaptive mesh
- Discontinuity singularity after 6.3 days

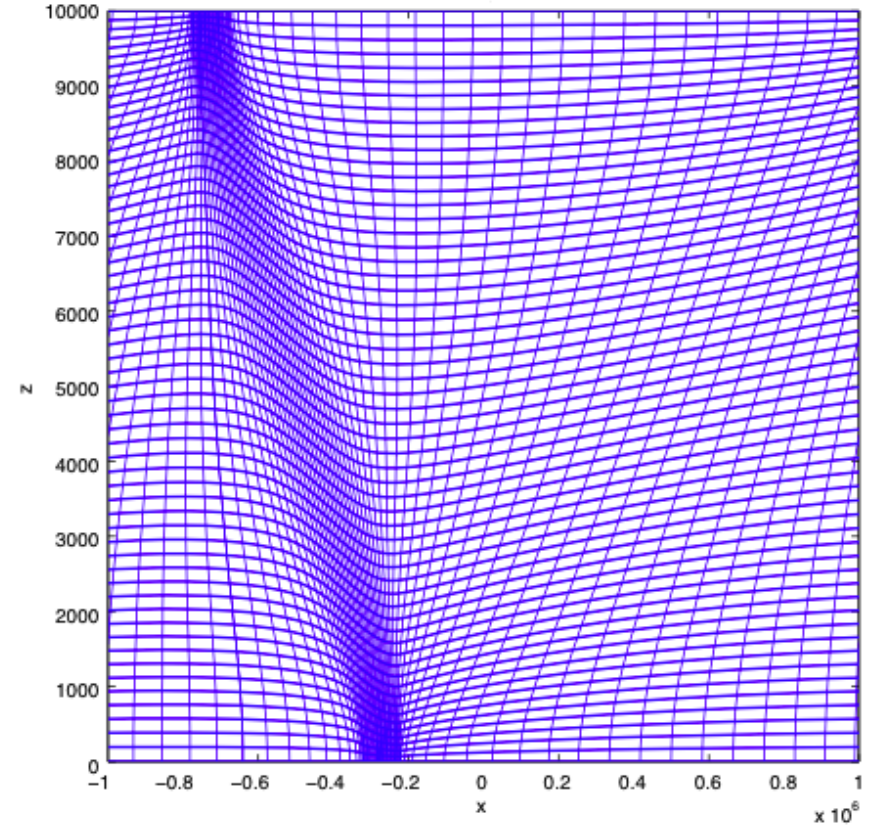
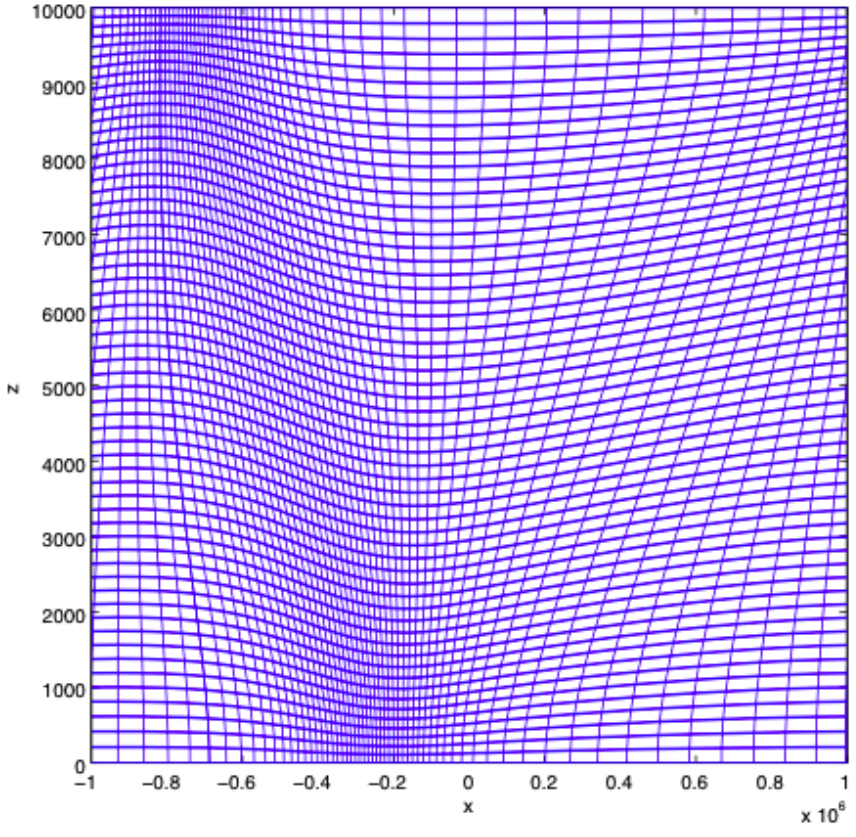


t = 5 days





t = 6.3 days



Conclusions

- Optimal transport is a natural way to determine meshes in dimensions greater than one
- It can be implemented using a relaxation process by using the PMA algorithm
- Method works well for a variety of problems, and there are rigorous estimates about its behaviour
- Looking good on meteorological problems
- Still lots of work to be done eg. Finding efficient ways to couple PMA to the underlying PDE